

庐枞矿集区与钨矿床有关的花岗岩的年代学及地球化学特征: 岩石成因及其对长江中下游晚白垩世成矿的启示*

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Abstract The Middle and Lower Reaches of Yangtze River Metallogenic Belt is an important Cu–Au–Mo–Fe metallogenic belt in eastern China, most deposits in the belt are associated with Mesozoic magmatism. A new type of tungsten–molybdenum polymetallic deposit which is discovered in the Donggushan district, north of Luzong orefield, and it has a close relationship with the concealed granites. The petrological and geochemical studies indicated that the granitic rocks contained relatively high SiO₂, Na₂O and K₂O, low FeO and MgO, and are nearly saturated in Al₂O₃, suggesting that they are high-K calc-alkaline series rocks. These granites are also characterized by enrichment of highly incompatible elements such as Rb, Th and U as well as LREE, depletion of high field strong elements such as Nb, P, Ti and Y, as well as low $\epsilon_{Nd}(t)$ (–18.2 ~ –17.1) and relatively high (⁸⁷Sr/⁸⁶Sr)_i (0.70596 ~ 0.70631). It can be concluded that their parental magmas have been derived from partial melting of Yangtze lower crust. And the magmas had been undergone fractional crystallization before their emplacements. Zircon U–Pb LA–ICP–MS dating of the granites have yielded age of 96.7 ± 1.3 Ma, indicating that the granitic intrusive rocks formed at earlier stage of Late Cretaceous. Compared with the previous geochronology data, a new diagenesis–mineralization event in the Middle and Lower Reaches of Yangtze River Metallogenic Belt was identified by the concealed granitic magmatism and tungsten–molybdenum polymetallic mineralization.

Key words Granite; Zircon U–Pb dating; Petrogeochemistry; Middle and Lower Reaches of Yangtze River Metallogenic Belt; Donggushan district, Anhui

摘要 长江中下游是中国东部一个重要的与中生代岩浆作用有关的 Cu–Au–Mo–Fe 矿成矿带。近期勘查工作在庐枞矿集区东顾山地区深部首次发现了与隐伏花岗岩密切相关的钨钼多金属矿化。岩石学和岩石地球化学研究结果表明, 隐伏花岗岩具有高硅富碱高钾、贫铁贫镁及铝弱饱和的特征, 属于高钾钙碱性系列岩石; 具有富集 Rb、Th、U 等强不相容元素和轻稀土

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元素,亏损 Nb、P、Ti、Y 等高场强元素,低 $\epsilon_{Nd}(t)$ ($-18.2 \sim -17.1$) 和相对高的 ($^{87}\text{Sr}/^{86}\text{Sr}$)_i (0.70596 ~ 0.70631) 的特征,推测其岩浆来源于扬子下地壳的部分熔融,岩浆在就位之前经历了一定程度的分异结晶作用。利用 LA-ICP MS 锆石 U-Pb 定年获得 $96.7 \pm 1.3\text{Ma}$ 的锆石年龄,表明其属于晚白垩世早期岩浆活动的产物。结合区域成岩成矿时空格架,认为东顾山地区隐伏花岗岩及与之有关的钨钼多金属矿化可能代表了长江中下游成矿带的新一期成岩成矿事件。

关键词 花岗岩; 锆石 U-Pb 定年; 岩石地球化学; 长江中下游成矿带; 安徽东顾山地区

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长江中下游成矿带是我国重要的 Cu-Au-Mo-Fe 矿成矿带(常印佛等,1991; 翟裕生等,1992; 唐永成等,1998; 毛景文等,2004; 周涛发等,2008)。成矿带中生代岩浆活动强烈,具有爆发性、阶段性和分区性的特征(常印佛等,1991; 翟裕生等,1992; Pan and Dong, 1999)。自二十世纪八十年代以来,前人对成矿带的岩浆岩进行了大量的科学研究,在岩石类型、成岩物质来源、成岩构造背景、岩浆岩成矿专属性等方面取得了许多重要的科学成果(常印佛等,1991; 翟裕生等,1992; 唐永成等,1998; 华仁民和毛景文,1999; 王强等,2003; 杜建国等,2003; 周涛发等,2008; 孙卫东等,2010; 毛景文等,2012) 特别是随着同位素定年技术的飞跃进步和对长江中下游地区成岩时代研究的不断深入,获得了一系列高质量的岩浆岩同位素年龄数据,建立了成矿带 148 ~ 135Ma、135 ~ 127Ma、127 ~ 121Ma 和 109 ~ 102Ma 四个阶段的成岩时空格架(周涛发等,2008; 曾键年等,2013; 孙洋等,2014; 刘建敏等,2014)。近年来,长江中下游成矿带由于深部找矿的突破而更加备受关注(常印佛等,1991; 董树文等,2007, 2010; 吕庆田等,2004, 2014; Li *et al.*, 2013; 吴明安等,2007; 张舒等,2014; 熊欣等,2014)。2014 年底,在庐枞矿集区北部深部勘探过程中,发现了东顾山钨钼多金属矿化及与成矿有关的隐伏花岗岩侵入体,这一发现不仅填补了庐枞矿集区乃至长江深断裂带以北地区无钨矿的空白,而且为研究长江中下游成矿带的岩浆作用提供了新的方向。聂利青等(2016) 仅对东顾山钨钼多金属矿床进行了成岩成矿年龄的测定,并进一步完善了长江中下游成矿带的成岩成矿时空格架。但对于与东顾山钨钼多金属矿床有关的花岗岩的岩石地球化学特征、岩石成因及与成矿带中不同阶段岩浆岩的对比研究工作至今还未开展。本文在详细的野外地质调查和室内岩相学观察的基础上,系统测定了与钨钼多金属矿床有关的花岗岩的化学成分和 Sr、Nd 同位素组成,并与断隆区、断凹区、宁镇地区岩浆岩及沿江 A 型花岗岩进行对比,探讨了它们在岩石种属、岩浆源区及成岩构造背景等方面的差异; 利用 LA-ICP MS 锆石 U-Pb 法进行了花岗岩的成岩时代的测定,结合长江中下游成矿带已有岩浆岩的年代学数据,重新构建了成矿带成岩时空格架,进一步阐述了东顾山地区燕山期岩浆-成矿作用可能为长江中下游成矿带的新一期成岩成矿作用。

1 地质背景

长江中下游成矿带位于扬子板块北缘的长江断裂带内,

大别-苏鲁晚三叠世碰撞造山带的前陆,经历了中元古界基底(董岭岩群)形成阶段、古生界盖层沉积阶段和中生代板内变形阶段(常印佛等,1991),长期的构造作用、岩浆活动和成矿作用使其形成了断隆和断凹的构造格架(周涛发等,2008)(图 1)。成矿带由七个各具特点的矿集区组成,自西向东依次为鄂东南、九瑞、安庆-贵池、铜陵、庐枞、宁芜及宁镇矿集区。庐枞矿集区位于长江中下游成矿带中部,包括庐枞火山岩盆地、沙溪地区和东顾山地区。

东顾山地区位于庐枞矿集区北部,区内不同程度出露震旦纪-白垩纪地层,震旦纪地层为一套海相沉积的碳酸盐岩; 寒武纪-晚二叠世地层为本区稳定发育地沉积盖层,总体为一陆表海环境,形成一套碳酸盐岩、陆源碎屑岩和硅质岩组合; 侏罗纪-白垩纪地层为一套陆相碎屑岩沉积。赋矿地层为早奥陶世红花园组浅灰色中-厚层状含硅质砂页岩夹泥质灰岩、生物碎屑灰岩,在地表发育强烈的硅化。东顾山地区构造较为复杂,褶皱构造主要有龙池山-东顾山向斜和盛桥-照壁山背斜,均属于斜歪-倒转褶皱类型,矿区位于盛桥-照壁山复式背斜的西部。断裂构造以 NE 向走向断裂和 NW 向横断裂为主,其中 NW 向断裂为一系列左旋走滑平移断层,错断早期形成的 NE 向走向断裂,较大规模的有东顾山和二蛟子山-照壁山走向断裂。矿区岩浆岩较为发育,主要包括东顾山隐伏花岗岩岩体和地表出露的冶父山岩体、吕家院岩体及少量的闪长玢岩脉等(图 2a)。东顾山隐伏花岗岩与钨钼多金属矿化直接相关,为成矿岩体; 冶父山岩体为石英正长岩,以小岩株状产出,呈不规则长条状 NNE 向展布,出露面积约 4.6km²,岩体接触带及内部含有大量的含金石英脉(邱检生和朱成伟,1999); 吕家院岩体为二长斑岩,零星分布在冶父山岩体周围及以捕虏体形式分布在冶父山岩体内部,出露面积约 0.35km²,岩体深部可能为一规模较大的岩株,其中心部位被后期的冶父山岩体吞蚀。

经深部钻孔揭露发现,东顾山地区隐伏花岗岩体顶部发育多个钨矿化体(图 3e, f),主要以脉状矿化为主,次为浸染状矿化; 在花岗岩体与早奥陶世红花园组碳酸盐岩地层接触带外带的石榴子石砂卡岩中发现钨钼共生矿化体,矿体形态以似层状、平缓透镜状为主,以浸染状矿化为主; 在外接触带的透辉石、透闪石砂卡岩带中及早奥陶世五峰组硅质岩、志留系碎屑岩与奥陶系碳酸盐岩界面-硅钙面中发现多处砂卡岩型铜、铅、锌、银多金属矿化体,主要由平行缓倾斜(0° ~ 20°)粗脉和不同产状的网脉组成(图 2b)。

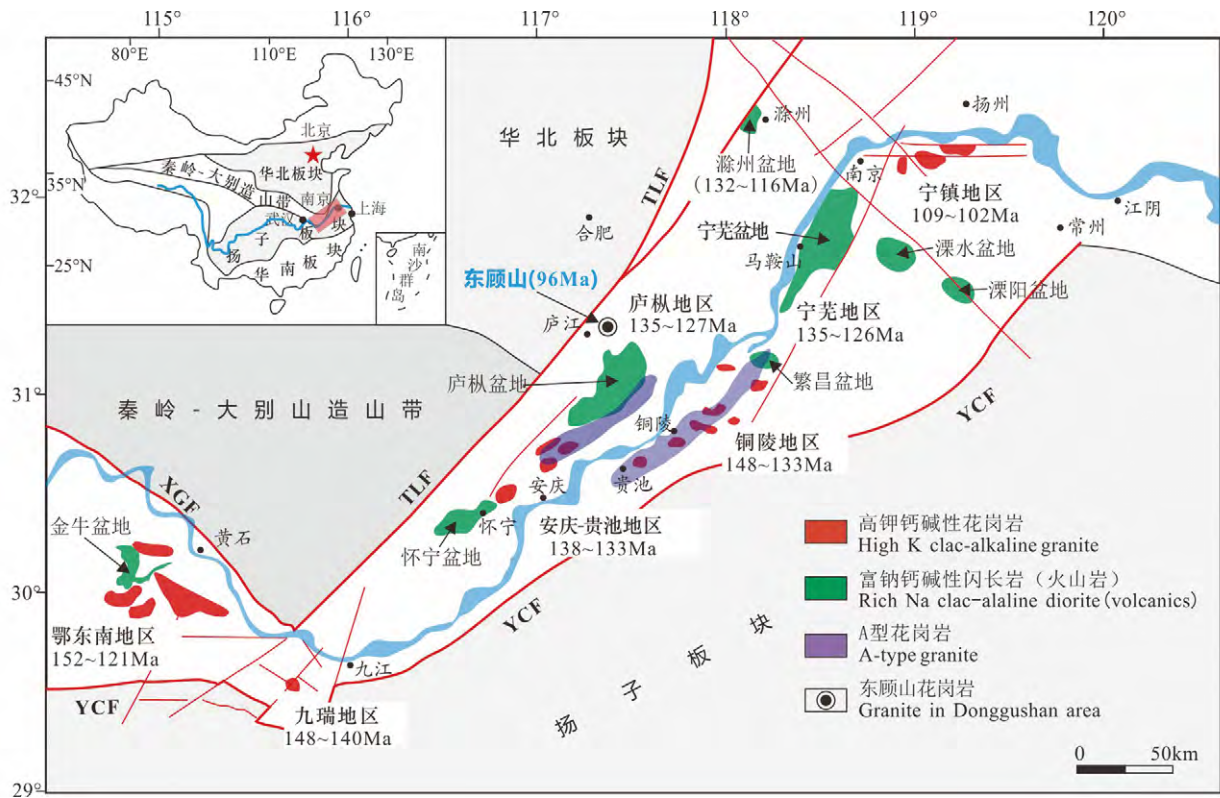


图1 长江中下游成矿带主要岩浆岩分布略图(据毛景文等, 2012 修改)

Fig.1 Sketch magma distribution map of the Middle and Lower Reaches of Yangtze River Metallogenic Belt (modified after Mao *et al.*, 2012)

2 样品特征

本文选取东顾山隐伏岩体新鲜的花岗岩样品开展岩石化学分析、Sr-Nd 同位素组成分析和锆石 U-Pb 定年测试工作, 实验样品地质特征如下:

新鲜实验样品采自矿区已施工钻孔 ZK002 中(深度范围为 1140 ~ 1200m), 岩性为花岗岩, 岩石呈灰白色, 半自形不等粒结构, 块状构造(图 3a), 主要矿物组成为石英(25%)、钾长石(40%)、斜长石(30%)、黑云母(<5%)及少量副矿物。其中石英呈灰白色, 部分为无色透明, 呈他形粒状分布在长石颗粒之间, 粒径大小约 0.5 ~ 2mm; 钾长石呈灰白色, 以半自形短柱状为主, 部分为不规则粒状, 大小为 0.5 × 0.5mm ~ 2 × 3mm, 部分短柱状钾长石颗粒中包裹有较早形成的长板状斜长石, 斜长石具环带, 大小为 0.2 × 0.5mm 左右, 钾长石表面多为弥漫状; 斜长石以灰白色为主, 多为长板状, 少量为不规则粒状, 大小为 0.2 × 0.5mm ~ 1.5 × 3mm, 部分斜长石边部见一圈钾长石环带; 黑云母呈褐色, 片状, 解理清晰, 大小不等, 约 0.2 × 0.6mm ~ 1 × 1.5mm, 主要以片状和片状集合体分布在长石颗粒之间(图 3b)。在侵入体边部为花岗斑岩(图 3c), 斑晶以钾长石和石英为主, 基质为细粒石英和长石(图 3d), 岩石中见黑云母二长花岗岩包体。

3 分析方法

3.1 全岩主微量元素测试

全岩主微量元素测试分析由核工业北京地质研究院分析测试研究中心完成。将新鲜的岩石样品粉碎至 200 目以下的粉末。主量元素分析在飞利浦 PW2404 X 射线荧光光谱仪上完成, 采用 GB/T 14506.28-93 硅酸盐岩石化学分析方法 X 射线荧光光谱法(XRF)分析测定, 相对误差小于 5%; 包括稀土元素在内的微量元素采用 Finnigan MAT 制造的 HR-ICP-MS(Element I) 仪器上完成, 实验方法采用电感耦合等离子体质谱(ICP-MS)方法通则。实验过程中温度为 20℃, 相对湿度 30%; 微量元素含量大于 10×10^{-6} 时, 相对误差 < 5%, 含量小于 10×10^{-6} 时, 相对误差 < 10%, 实验详细流程参见靳新娣和朱和平(2000)。

3.2 全岩 Sr-Nd 同位素测试

全岩 Sr-Nd 同位素组成分析在核工业北京地质研究所分析测试研究中心完成。Rb-Sr 同位素分析采用 ISOPROBE-T 热电质谱计, 单带, M+, 可调多法拉第接收器接收; 质量分馏用 $^{86}\text{Sr}/^{88}\text{Sr} = 0.1194$ 校正, 标准测量结果: NBS987 =

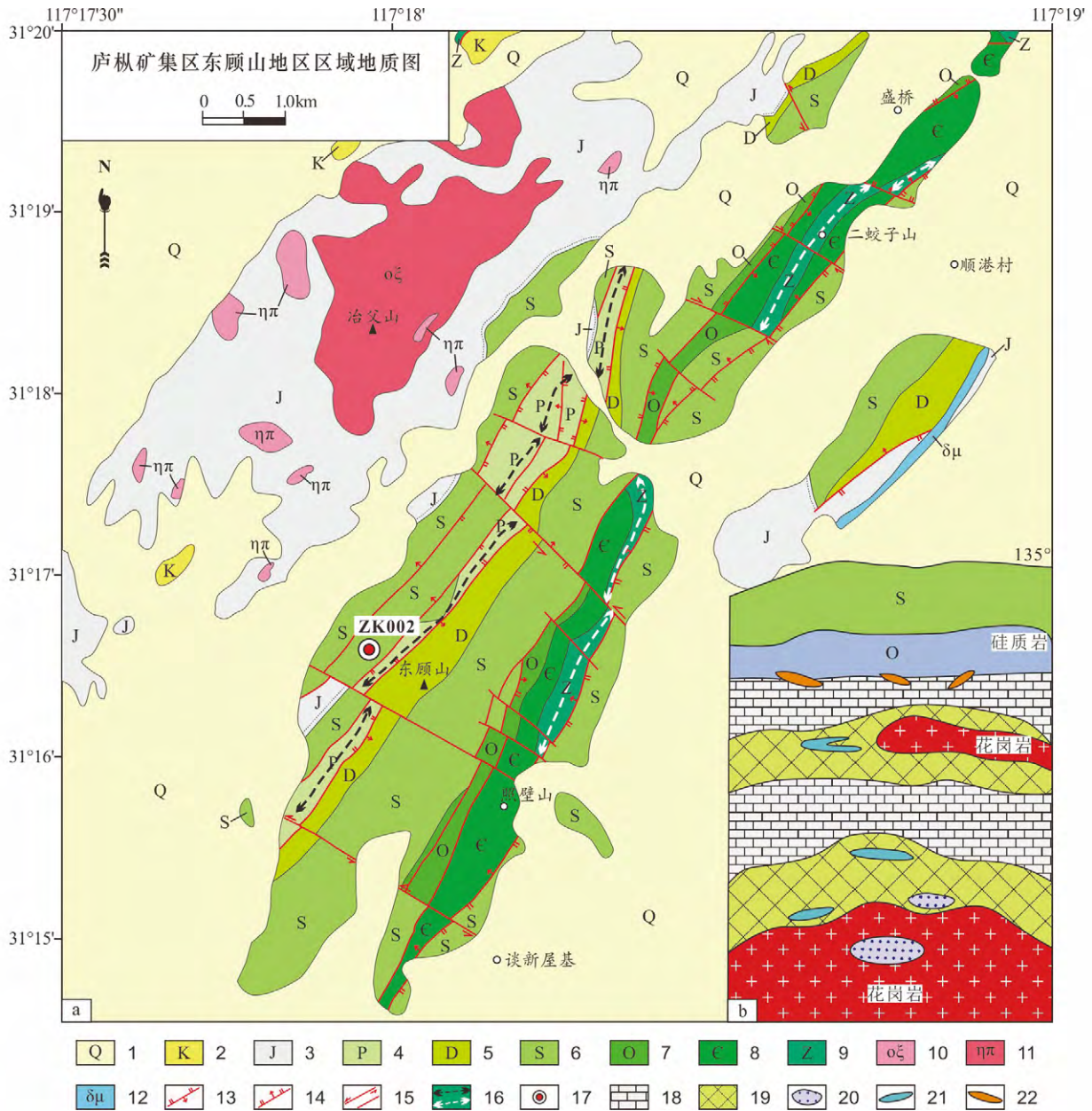


图2 庐枞矿集区东顾山地区区域地质图(a)及矿体产出示意图(b) (据聂利青等, 2016 修改)

1-第四系; 2-白垩系; 3-侏罗系; 4-二叠系; 5-泥盆系; 6-志留系; 7-奥陶系; 8-寒武系; 9-震旦系; 10-石英正长岩; 11-二长斑岩; 12-闪长玢岩; 13-逆断层; 14-正断层; 15-平移断层/性质不明断层; 16-向斜轴/背斜轴; 17-钻孔; 18-大理岩; 19-砂卡岩; 20-钼矿体; 21-钨矿体; 22-铅锌矿体

Fig. 2 Regional geological map of Donggushan district in Luzong orefield (a) and a cross section of the Donggushan tungsten polymetallic deposit (b) (modified after Nie *et al.*, 2016)

1-Quaternary; 2-Cretaceous; 3-Jurassic; 4-Permian; 5-Devonian; 6-Silurian; 7-Ordovician; 8-Cambrian; 9-Sinian; 10-quartz syenite; 11-monzonite porphyry; 12-dioritic porphyrite; 13-reverse fault; 14-normal faults; 15-Strike-slip fault and supposed fault; 16-synclinal axis and anticlinal axis; 17-drill; 18-marble; 19-skarn; 20-molybdenum orebody; 21-tungsten orebody; 22-lead-zinc orebody

0.710250 ± 7, 实验室流程本底: Rb 为 2×10^{-10} g, Sr 为 2×10^{-10} g。Sm-Nd 同位素分析采用 ISOPROBE-T 热电质谱计, 三带, M+, 可调多法拉第接收器接收; 质量分馏用 $^{146}\text{Nd}/^{144}\text{Nd} = 0.7219$ 校正 标准测量结果: JMC 为 $^{143}\text{Nd}/^{144}\text{Nd} = 0.512109 \pm 3$, 全流程本底 Sm 和 Nd 均小于 50pg。相关的化学流程和同位素比值测试可参见 Chen *et al.* (2000)。

3.3 锆石 U-Pb 年龄测试

定年样品的锆石挑选工作由河北省廊坊诚信地质服务公司完成。在详细地野外地质观察及岩石学研究的基础上, 将 3~5kg 的原岩样品粉碎, 采用标准重矿物分离技术, 经过淘洗、重选、磁选和密度分选等流程, 在双目镜下进一步挑选

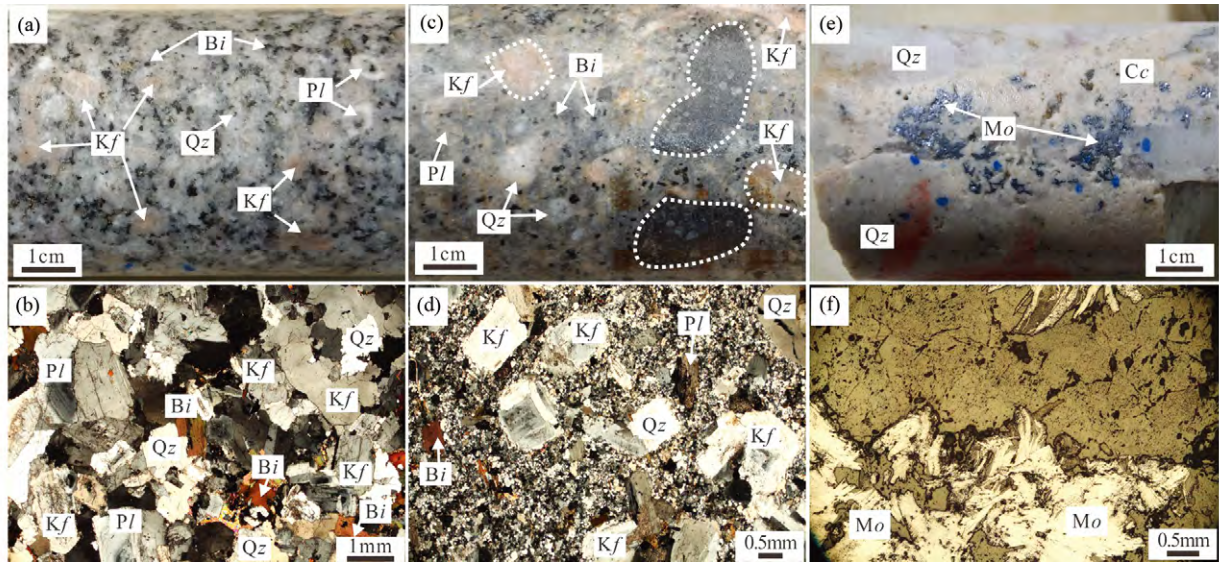


图3 东顾山隐伏花岗岩岩石及矿化样品的手标本和显微照片

(a) 隐伏花岗岩手标本; (b) 隐伏花岗岩显微照片(正交偏光) 花岗结构; (c) 隐伏花岗岩斑岩手标本, 斑晶为钾长石和石英, 岩石中含有黑云母二长花岗岩包体(虚线内); (d) 隐伏花岗岩斑岩显微照片(正交偏光) 斑晶为钾长石和石英, 少量为黑云母, 基质为石英和长石; (e) 岩体中石英-方解石-辉钼矿脉, 辉钼矿呈鳞片状集合体, 位于脉中间; (f) 辉钼矿脉显微照片(反射光)。Pl-斜长石; Bi-黑云母; Kf-钾长石; Qz-石英; Cc-方解石; Mo-辉钼矿

Fig. 3 Photomicrographs of concealed granite and mineralization rocks samples in Donggushan area

(a) concealed granite; (b) micrograph of concealed granite (CPL) with granitic texture; (c) concealed granite porphyry with the phenocrysts being potassium feldspar and quartz contains some biotite adamellite enclaves (within the dotted line); (d) micrograph of concealed granite porphyry (CPL) with the phenocrysts being potassium feldspar, quartz and less biotite, and the matrix being quartz and feldspar; (e) quartz-calcite-molybdenite vein in granite intrusive rocks, molybdenite occurred in schistose aggregated; (f) micrograph of molybdenite vein (reflected light). Pl-plagioclase; Bi-biotite; Kf-K-feldspar; Qz-quartz; Cc-calcite; Mo-molybdenite

表1 东顾山隐伏花岗岩主量元素分析结果表(wt%)

Table 1 The major elements (wt%) analytical results of concealed granite intrusions in Donggushan area

样品号	SiO ₂	TiO ₂	Al ₂ O ₃	Fe ₂ O ₃	FeO	MnO	MgO	CaO	Na ₂ O	K ₂ O	P ₂ O ₅	LOI	Total	Na ₂ O + K ₂ O	K ₂ O / Na ₂ O	ANK	ANCK	σ
DGSZK002-3	70.38	0.41	14.97	2.65	1.30	0.04	0.95	2.25	4.28	3.90	0.17	1.17	98.80	8.18	0.91	1.35	0.97	2.44
DGSZK002-4	69.31	0.50	15.35	2.80	1.26	0.04	0.98	2.35	4.32	4.16	0.19	1.27	98.70	8.48	0.96	1.35	0.97	2.73
DGSZK002-7	73.62	0.30	13.65	1.95	1.23	0.03	0.68	1.49	3.83	4.34	0.11	0.93	99.00	8.17	1.13	1.26	0.99	2.18
DGSZK002-8	75.46	0.23	13.01	1.47	0.69	0.03	0.42	1.12	3.67	4.51	0.07	0.86	99.00	8.19	1.23	1.21	1.00	2.06

注: ANK = Al₂O₃ / (K₂O + Na₂O) 原子量比值; ANCK = Al₂O₃ / (K₂O + Na₂O + CaO) 原子量比值; σ = (K₂O + Na₂O)² / (SiO₂ - 43)

出晶形完好、透明度和色泽较好的锆石颗粒。将分选后的锆石颗粒粘贴于环氧树脂靶之上, 打磨刨光, 然后进行锆石的透反射光、阴极发光拍照, 根据锆石的形态特征, 确定激光测年的位置。

LA-ICP-MS 锆石 U-Pb 测年由西北大学大陆动力学国家重点实验室完成。所使用的 ICP-MS 为 Elan 6100DRC, 激光剥蚀系统为 GeoLas 200M 深紫外(DUV) 193nm ArF 准分子激光剥蚀系统。分析所采用的光斑直径为 32μm, 并采用 ²⁹Si 作为内标, 以哈佛大学标准锆石 91500 作为外标校正。锆石 U-Pb 年代学测试数据的处理采用 Glitter(4.0 版) 软件进行, 普通铅校正采用 Andersen 的方法 (Andersen, 2002), 年龄计算采用 Isoplot(3.23 版) 进行 (Ludwing, 2003), 测试中的误差

标准为 1σ, 实验详细的流程参见袁洪林等 (2003) 和 Yuan *et al.* (2004)。本次实验有效的测试数据为 14 个。

4 测试结果

4.1 主量元素

表 1 列出了东顾山隐伏花岗岩的主量元素分析结果。从表中看出, SiO₂ 含量变化范围为 69.31% ~ 75.46%, K₂O 含量为 3.90% ~ 4.51%, 全碱含量介于 8.17% ~ 8.48% 之间, 高于中国同酸度岩石的全碱平均含量 (黎彤等, 1998), 具有富碱特征。在侵入岩硅碱图 (图 4) 中, 东顾山隐伏侵入岩样品全部落于花岗岩区域内。在岩石系列划分上, 本区隐

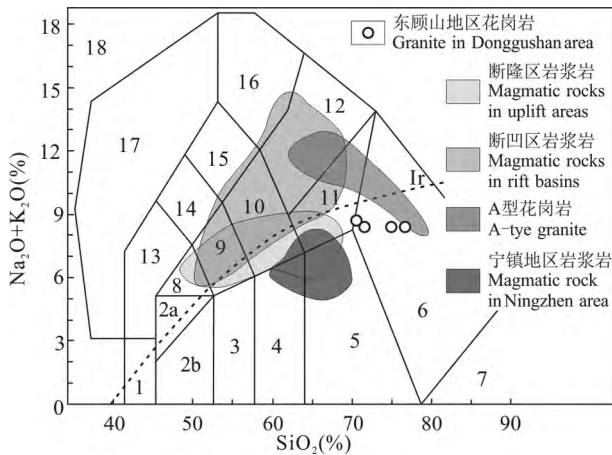


图4 东顾山隐伏花岗岩 SiO_2 -($\text{K}_2\text{O} + \text{Na}_2\text{O}$) 图解(底图据 Middlemost, 1994)

图4-图8 阴影区的数据来源:断隆区(陈江峰等,1993;王强等,2003,2004;薛怀民等,2006;于文修,2006;李进文等,2007;蒋少涌等,2008;吴才来等,2010;Mao *et al.*,2011;徐晓春等,2012;谢建成等,2012);断凹区(任启江等,1991;刘洪等,2002;谢智等,2007;谢成龙等,2007;邓晋福等,2011;袁峰等,2011;薛怀民等,2012,2016;范裕等,2014;张舒等,2014);A型花岗岩(楼亚儿和杜杨松,2006;范裕等,2008;曹毅等,2008;向文帅等,2009;Li *et al.*,2012;彭戈等,2012);宁镇地区(宁仁祖和陈根生,1989;许继峰等,2001;徐莺,2010;洪文涛等,2010)

Fig. 4 SiO_2 vs. ($\text{K}_2\text{O} + \text{Na}_2\text{O}$) diagram of concealed granite in Donggushan area (after Middlemost, 1994)

Data sources of the shadow region in Fig. 4 to Fig. 8: Uplift areas (Chen *et al.*, 1993; Wang *et al.*, 2003, 2004; Xue *et al.*, 2006; Yu, 2006; Li *et al.*, 2007; Jiang *et al.*, 2008; Wu *et al.*, 2010; Mao *et al.*, 2011; Xu *et al.*, 2012; Xie *et al.*, 2012); Rift basins (Ren *et al.*, 1991; Liu *et al.*, 2002; Xie *et al.*, 2007; Xie *et al.*, 2007; Deng *et al.*, 2011; Yuan *et al.*, 2011; Xue *et al.*, 2012, 2016; Fan *et al.*, 2014; Zhang *et al.*, 2014); A-type granite (Lou and Du, 2006; Fan *et al.*, 2008; Cao *et al.*, 2008; Xiang *et al.*, 2009; Li *et al.*, 2012; Peng *et al.*, 2012); Ningzhen area (Ning and Chen, 1989; Xu *et al.*, 2001; Xu, 2010; Hong *et al.*, 2010)

伏花岗岩与宁镇地区岩浆岩一致,均为亚碱性系列岩石,而与断凹区岩浆岩截然不同,与断隆区岩浆岩和沿江A型花岗岩存在一定的差异。进一步将亚碱性系列岩石进行划分,东顾山隐伏花岗岩全部落入钙碱性系列范围内,相对于宁镇地区岩浆岩和断隆区岩浆岩,本区隐伏花岗岩更靠近AF一侧,远离FM一侧(图5),具有更富碱贫铁贫镁的特征。在硅钾图上(图6),本区隐伏花岗岩全部落入高钾钙碱性系列范围,与断隆区岩浆岩和宁镇地区岩浆岩一致,其中断隆区部分样品分布在钾玄岩系列范围的原因可能与岩石蚀变(钾化)有关(徐晓春等,2012;谢建成等,2012)断凹区岩浆岩和沿江A型花岗岩富钾,全部落在钾玄岩系列范围内。本区花岗岩 $\text{ANCK} = 0.97 \sim 1.00$,且岩石矿物组合中见黑云母,表明东顾山隐伏花岗岩具弱过铝质特征。

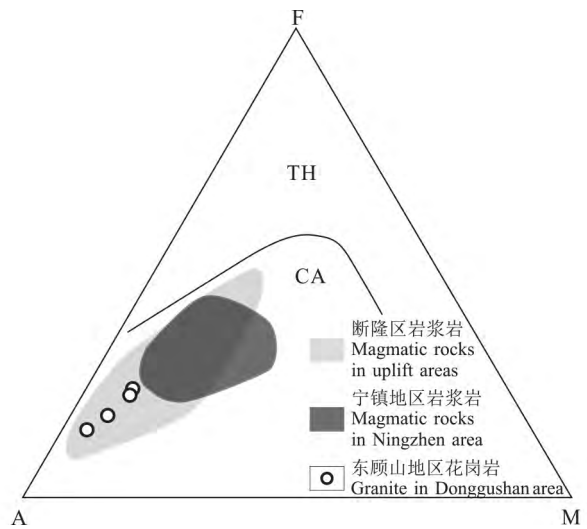


图5 东顾山隐伏花岗岩 AFM 判别图解(底图据 Irvine and Baragar, 1971)

Fig. 5 AFM diagram of concealed granite in Donggushan area (after Irvine and Baragar, 1971)

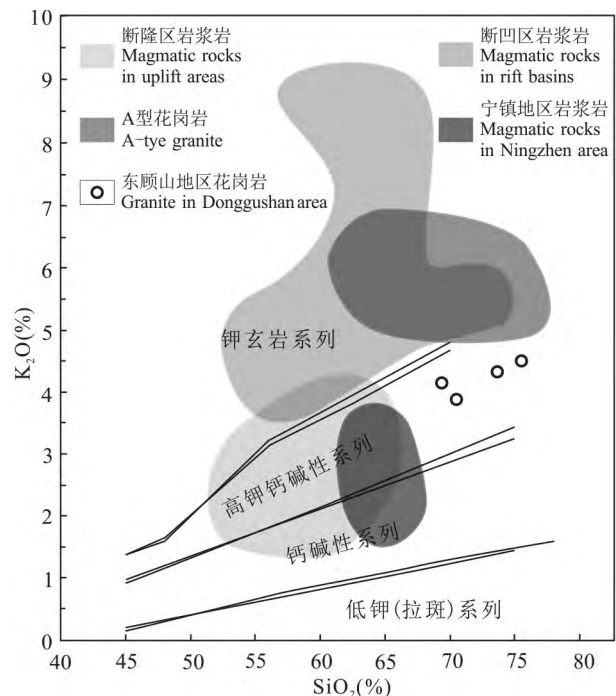


图6 东顾山隐伏花岗岩 SiO_2 - K_2O 判别图解(底图据 Middlemost, 1994)

Fig. 6 SiO_2 vs. K_2O diagram of concealed granite in Donggushan area (after Middlemost, 1994)

4.2 稀土元素

东顾山隐伏花岗岩的稀土元素分析结果见表2,稀土元素总量介于 $147.5 \times 10^{-6} \sim 252.8 \times 10^{-6}$ 之间,平均值为 190.3×10^{-6} ,低于世界花岗岩质岩石的稀土元素平均含量

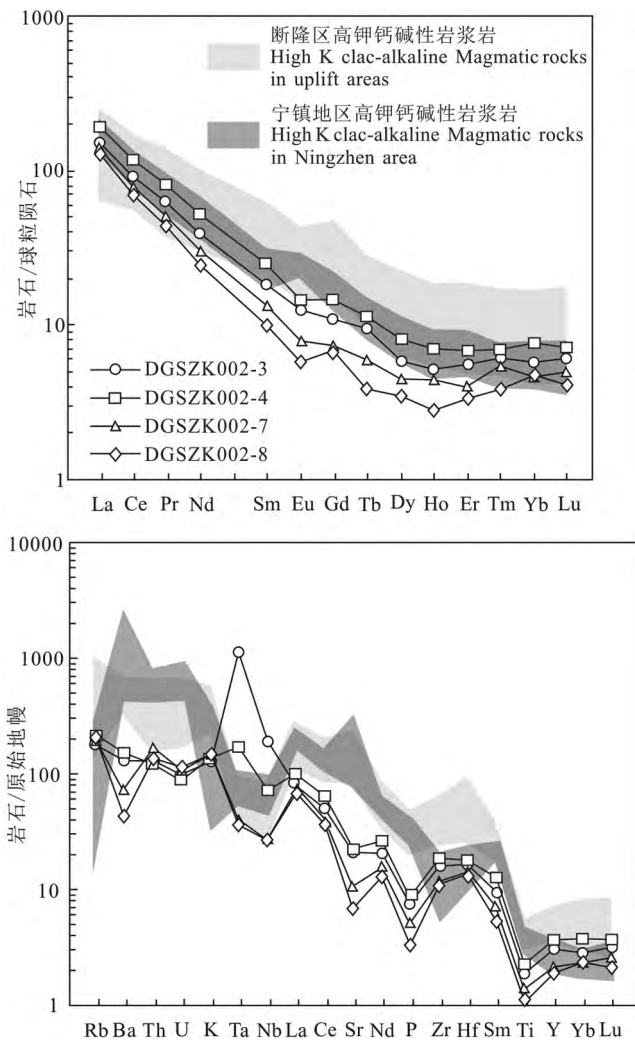


图7 东顾山隐伏花岗岩的球粒陨石标准化稀土元素配分图和原始地幔标准化微量元素蛛网图(标准化值据 Sun and McDonough , 1989)

Fig.7 Chondrite-normalized rare earth elements patterns and primitive mantle-normalized trace element patterns for the concealed granite in Donggushan area (normalization values after Sun and McDonough , 1989)

(290×10^{-6} , Taylor and Sheppard, 1986)。本区隐伏花岗岩轻重稀土元素分异明显, $(La/Yb)_N$ 介于 25.1 ~ 30.2, 平均值为 26.9。球粒陨石标准化稀土元素配分模式显示(图7), 本区侵入岩的稀土元素配分模式与断隆区和宁镇地区高钾钙碱性系列岩浆岩的稀土元素配分曲线形式较为相似, 均为向右倾斜的轻稀土元素富集型, 但在重稀土元素含量上, 与两者存在一定的差异。东顾山隐伏花岗岩 δEu 变化范围为 0.67 ~ 0.86, 平均值为 0.76, 表现为 Eu 负异常, 与断隆区高钾钙碱性系列岩浆岩基本一致, 而宁镇地区岩浆岩具有弱 Eu 正异常, 可能由于发生过斜长石的堆晶作用引起。

表2 东顾山隐伏花岗岩微量元素分析结果表 ($\times 10^{-6}$)

Table 2 The trace elements ($\times 10^{-6}$) analytical results of concealed granite intrusions in Donggushan area

样品号	Li	Be	Se	V	Cr	Co	Ni	Cu	Zn	Ga	Rb	Sr	Mo	Cd	In	Sb	Cs	Ba	W	Re	Tl
DGSZK002-3	25.4	3.59	4.34	35.7	7.71	5.15	5.42	116	26.5	17.5	115	456	0.344	0.148	0.014	0.029	1.06	934	4.68	0	0.498
DGSZK002-4	31.4	3.21	4.87	39.9	5.96	4.62	5.42	45.3	24.3	19.9	139	460	0.388	0.366	0.007	0.022	1.52	1103	3.68	0.01	0.509
DGSZK002-7	14.9	2.03	2.6	24.6	4.55	2.28	3.99	18.2	24.7	16.6	130	229	0.103	0.173	0.011	0.039	0.739	499	0.303	0.01	0.338
DGSZK002-8	6.71	2.27	2.29	18.2	10.8	1.94	5.54	8.54	21.8	14.9	131	146	0.717	0.122	0.004	0.057	1.06	309	0.189	0.01	0.444
样品号	Pb	Bi	Th	U	Nb	Ta	Zr	Hf	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb
DGSZK002-3	15.6	0.178	11.1	2.13	136	46.9	184	5.11	54	89	8.66	28.4	4.28	1.11	3.38	0.554	2.31	0.444	1.42	0.221	1.45
DGSZK002-4	14.5	0.064	10.4	1.86	51.8	7.03	209	5.59	70.3	114	11.4	37	5.77	1.27	4.52	0.663	3.1	0.601	1.71	0.252	1.89
DGSZK002-7	17.5	0.016	14.7	2.32	19.5	1.68	135	4.5	51.4	74.6	6.96	21.6	3.13	0.691	2.26	0.359	1.74	0.388	1.01	0.204	1.15
DGSZK002-8	20.4	0.028	11.7	2.48	19.6	1.53	128	4.3	47.8	67.1	6.07	17.4	2.35	0.499	2.11	0.23	1.32	0.246	0.85	0.138	1.19
样品号	Lu	Y	La/Nb	K/Nb	Ba/Nb	Rb/Sr	Sr/Y	Sr/Yb	La/Yb	Pb/Nd	Nb/Ta	Zr/Hf	(La/Yb) _N	δEu	ΣREE	LREE	HREE	LREE/HREE	LREE/HREE	LREE/HREE	LREE/HREE
DGSZK002-3	0.234	14.1	0.26	2.02	1.19	414	0.017	155	0.86	1.85	2.9	36.0	25.2	0.86	0.88	195.5	185.5	10.0	18.5	18.5	18.5
DGSZK002-4	0.28	17.1	0.22	1.98	0.97	348	0.018	248	0.88	1.86	7.4	37.4	25.1	0.73	0.87	252.8	239.7	13.0	18.4	18.4	18.4
DGSZK002-7	0.193	10.1	0.22	2.22	1.44	348	0.014	110	0.95	1.80	11.6	30.0	30.2	0.76	0.82	165.7	158.4	7.30	21.7	21.7	21.7
DGSZK002-8	0.157	8.73	0.21	2.58	1.83	520	0.014	77.2	0.97	1.78	12.8	29.8	27.1	0.67	0.80	147.5	141.2	6.24	22.6	22.6	22.6

4.3 微量元素

东顾山隐伏花岗岩的微量元素分析结果见表2。从表中可以看出该隐伏侵入体具有相对较高含量的 Rb、Th、U 等强不相容元素和轻稀土元素,相对较低含量的高场强元素(HFSE: Nb、P、Ti、Y),Ba、Sr 具有负异常,Ta 既表现为富集又表现为亏损。

Sr 易取代斜长石中的 Ca,主要富集在富钙斜长石、磷灰石等,高 Sr 是幔源金伯利岩、大陆碱性玄武岩和橄榄玄武岩等高钾岩石的特征(邢凤鸣和徐祥,1995,1996)。东顾山隐伏花岗岩贫 Sr,其值介于 $146 \times 10^{-6} \sim 460 \times 10^{-6}$,平均值为 322×10^{-6} ,说明其形成过程中,富钙斜长石、磷灰石等含 Sr 矿物经历了分离结晶作用,同时暗示本区隐伏花岗岩岩浆来源并非直接来自幔源。Rb、Ba 在含钾矿物中与 K 可以广泛地类质同象,主要载体矿物为钾长石和黑云母,但是元素在一定的物理化学条件下迁移和富集的性质不同,在结晶分异过程中,含 K 矿物的种类及数量都可影响 Rb、Ba 的分配,本区隐伏花岗岩富 Rb、贫 Ba,很可能是原始岩浆固有的特征。一般,幔源型或者壳幔混源同熔型花岗岩的 Ba、Sr 含量高,Rb 含量低,而壳型花岗岩类 Ba、Sr 含量低,且富 Rb(Fowler *et al.*, 2001, 2008; 谢建成等, 2012)。东顾山隐伏花岗岩贫 Ba、Sr,富 Rb 的特征,说明其可能为壳型花岗岩类。P、Ti 元素负异常是由于磷灰石和钛铁矿的进一步分离结晶,使 P、Ti 明显降低造成的。

Nb 和 Ta、Zr 和 Hf 具有相同的离子价态和相似的离子半径,因此它们具有相似的地球化学行为,特别是在与地幔演化相关的岩浆过程中很难发生分馏(Hofmann, 1988)。一般认为球粒陨石的 Nb/Ta 和 Zr/Hf 比值较高(Nb/Ta = 19.9 ± 0.6 , Zr/Hf = 34.3 ± 0.3) (Münker *et al.*, 2003), MORB 和 OIB 都具有相近的低于球粒陨石的 Nb/Ta 比值(前者约为 14.2, 后者约为 15.9 ± 0.6) (Münker *et al.*, 2003; Pfänder *et al.*, 2007) 和高于球粒陨石的 Zr/Hf 比值(34~42) (Büchl *et al.*, 2002; Pfänder *et al.*, 2007), 大陆地壳的平均 Nb/Ta 比值为 13.4 左右, Zr/Hf 比值为 36.7 左右(Rudnick and Gao, 2003)。研究表明,在大陆地壳分异和演化过程中, Nb/Ta 和 Zr/Hf 比值并非完全固定,彼此之间会发生分异(Dostal and Chatterjee, 2000; Hoffmann *et al.*, 2011), 引起分异的机制可能是相关矿物的分离结晶或者在高压下残留造成(Münker *et al.*, 2004), 也可能是流(熔)体与岩浆相互作用形成的(Green, 1994; Dostal and Chatterjee, 2000), 或者是在超临界流体媒介下热扩散引起 Nb 和 Ta 在低温部位发生重大分异(Ding *et al.*, 2009)。东顾山隐伏花岗岩的 Zr/Hf 比值介于 29.8~37.4, 变化范围小,基本接近大陆地壳的平均值,而 Nb/Ta 比值的变化范围为 2.9~12.8(其中前 2 个样品分别为 2.9 和 7.4, 后 2 个样品基本一致,且与大陆地壳的平均值一致),发生了较大的分异。Nb 和 Ta 的宿主矿物主要为含钛矿物,如角闪石、榍石、金红石、钛铁矿和含钛磁铁矿等。

东顾山隐伏花岗岩中黑云母的含量分布不均匀,前 2 个样品中黑云母含量远高于后 2 个样品,电子探针分析结果显示(未发表),黑云母中包裹着的大量磷灰石和含钛磁铁矿可能是引起 Nb、Ta、P、Sr 等元素含量相对较高的主要原因,而伴随亲铁元素 W、Cu 含量相对较高的原因是花岗岩中的黑云母及磁铁矿等富铁矿物是两者的主要载体。

4.4 Sr-Nd 同位素

东顾山隐伏花岗岩 Sr-Nd 同位素测试结果见表3。数据显示,本区花岗岩的 Sr 同位素初始值($^{87}\text{Sr}/^{86}\text{Sr}$)_i = 0.70596~0.70631, 平均 0.70610; Nd 初始值($^{143}\text{Nd}/^{144}\text{Nd}$)_i = 0.51158~0.51164, 平均 0.51161; $\varepsilon_{\text{Nd}}(t) = -18.2 \sim -17.1$, 平均 -17.7。

在 $\varepsilon_{\text{Nd}}(t) - (^{87}\text{Sr}/^{86}\text{Sr})_i$ 图解(图8)中,本区花岗岩样品集中分布于第四象限,靠近扬子下地壳,而远离扬子克拉通岩石圈地幔,意味着东顾山隐伏侵入岩的成岩物质主要来自扬子下地壳,可能含少量幔源端元物质的加入。断隆区侵入岩高钾钙碱性系列岩石在岩石圈地幔区域与扬子下地壳之间呈斜线排列,表明其物质来源是两种以上端元组成的混合物(邢凤鸣和徐祥, 1996; 陈江峰等, 1993; Wu *et al.*, 2000; 徐晓春等, 2012); 断凹区橄榄安粗岩系列岩石成岩物质来自富集型地幔,可能受到少量地壳物质的混染(杨荣勇等, 1993; 谢智等, 2007; 孙冶东等, 1994; 薛怀民等, 2010, 2016); A 型花岗岩分布在扬子克拉通岩石圈地幔与扬子上地壳之间,其成岩物质起源于交代地幔,在上侵过程中,受到地壳物质的混染(向文帅等, 2009; 张舒等, 2014; 薛怀民等, 2016)。可见,东顾山地区花岗岩在物质来源上与断隆区、断凹区及 A 型花岗岩存在很大的差异。

4.5 锆石 U-Pb 年龄

样品(DGSZK002-12)中锆石呈无色-淡黄色,多为半自形-自形柱状,偶见不规则状,锆石颗粒完整,少量锆石表面可见溶蚀现象,锆石长约 80~200 μm , 宽约 30~100 μm 。CL 图像显示(图9),锆石内部结构清晰,具有特征的单期生长的振荡环带,部分锆石颗粒颜色整体偏白,震荡环带表现为灰白-浅灰黑色交替出现,表明这些锆石中 U 的含量相对偏低,锆石中无继承锆石形成的核幔结构,无后期形成的变质壳,表明这些锆石由岩浆结晶形成。大量研究表明,不同成因锆石具有不同的 Th/U 比值,岩浆锆石 Th/U 比值一般大于 0.1(表4),而变质锆石 Th/U 比值一般小于 0.1(Belousova *et al.*, 2002; Watson *et al.*, 2006)。本文所测样品中锆石 Th/U 比值均大于 0.1,应属于典型的岩浆成因锆石,其测年结果可代表岩浆冷凝固结的时代。

锆石 U-Pb 定年结果表明,多数测点分布在 $^{206}\text{Pb}/^{238}\text{U} - ^{207}\text{Pb}/^{235}\text{U}$ 谐和图中谐和曲线附近(图10),谐和度均在 95% 以上,表明 U-Pb 体系封闭,未受后期热液事件干扰。东顾山隐伏花岗岩体(DGSZK002-12)的锆石表面年龄 $^{206}\text{Pb}/^{238}\text{U}$ 加

表3 东顾山地区隐伏花岗岩 Sr-Nd 同位素组成

Table 3 Sr-Nd isotopic compositions of concealed granite in Donggushan area

样品号	$\left(\frac{87\text{Rb}}{86\text{Sr}}\right)_m$	$\left(\frac{87\text{Sr}}{86\text{Sr}}\right)_m$	$\left(\frac{147\text{Sm}}{144\text{Nd}}\right)_m$	$\left(\frac{143\text{Nd}}{144\text{Nd}}\right)_m$	Age (Ma)	$\epsilon_{\text{Sr}}(0)$	$\epsilon_{\text{Sr}}(t)$	$f_{\text{Rb/Sr}}$	$\left(\frac{87\text{Sr}}{86\text{Sr}}\right)_i$	$\epsilon_{\text{Nd}}(0)$	$\epsilon_{\text{Nd}}(t)$	$f_{\text{Sm/Nd}}$	$t_{\text{DM}}(\text{Ma})$	$\left(\frac{143\text{Nd}}{144\text{Nd}}\right)_i$
DGSZK002-3	0.71233	0.707285	0.095615	0.511697		39.5	27.3	7.61	0.70631	-18.4	-17.1	-0.51	1873	0.511637
DGSZK002-4	0.853503	0.707123	0.09894	0.511662	96.7	37.2	22.3	9.32	0.70596	-19	-17.8	-0.50	1973	0.5116
DGSZK002-7	1.603453	0.708213	0.091937	0.511639		52.7	23.3	18.39	0.70603	-19.5	-18.2	-0.53	1889	0.511581
DGSZK002-8	2.534351	0.709574	0.085688	0.511658		72.0	24.6	29.65	0.70612	-19.1	-17.8	-0.56	1774	0.511604

注: 表中 Rb-Sr 和 Sm-Nd 同位素体系的计算公式一致, 现以 Sm-Nd 体系为例, 计算公式列举如下: $\epsilon_{\text{Nd}}(0) = \left[\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m / \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CHUR}} - 1 \right] \times 10^4$, 其中 $\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CHUR}} = 0.512638$; $\epsilon_{\text{Nd}}(t) = \left[\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m / \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CHUR}}(t) - 1 \right] \times 10^4$, 其中 $\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CHUR}}(t) = \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CHUR}} (e^{\lambda t} - 1) + \left(\frac{147\text{Sm}}{144\text{Nd}} \right)_{\text{CHUR}} / \lambda$; $f_{\text{Sm/Nd}} = \left[\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}} \right] / \left[\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}} - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CC}} \right]$; $t_{\text{DM}} = \frac{1}{\lambda} \ln \left[1 + \frac{\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}}}{\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}} - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CC}}} \right]$; 其中 $\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}} = 0.51315$, $\left(\frac{147\text{Sm}}{144\text{Nd}} \right)_{\text{DM}} = 0.2137$; $t_{\text{DM}} = 1/\lambda \ln \left[1 + \frac{\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}}}{\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{DM}} - \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CC}}} \right]$; 其中 $\left(\frac{147\text{Sm}}{144\text{Nd}} \right)_{\text{CC}} = 0.118$; $\left(\frac{143\text{Nd}}{144\text{Nd}} \right)_{\text{CC}} = \left(\frac{143\text{Nd}}{144\text{Nd}} \right)_m - \left(\frac{147\text{Sm}}{144\text{Nd}} \right)_m (e^{\lambda t} - 1)$, λ 为 ^{147}Sm 的衰变常数, 其值为 $6.54 \times 10^{-12} \text{y}^{-1}$

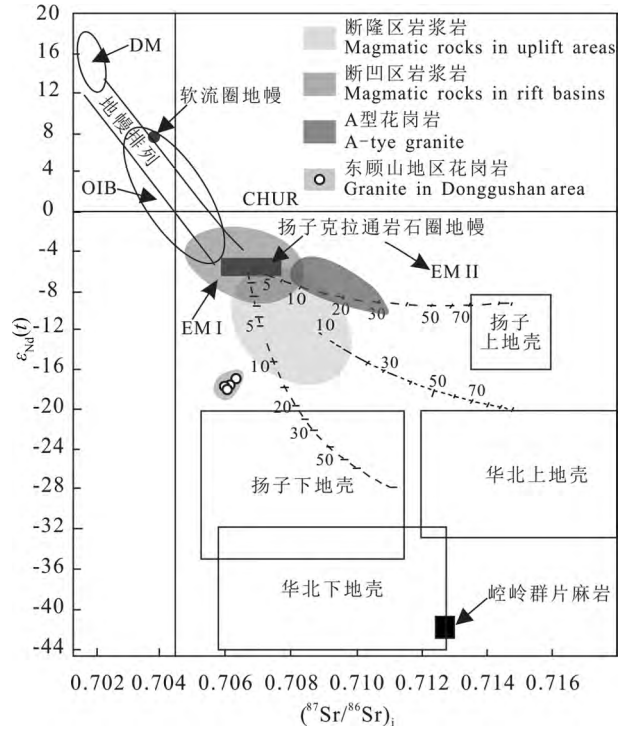


图8 东顾山隐伏花岗岩 Sr-Nd 同位素图解

数据来源: EM I 和 EM II 地幔端元 (Zindler and Hart, 1986); DM, OIB 和地幔序列 (Yang et al., 2005); 华北上地壳和下地壳 (Jahn et al., 1999); 扬子上地壳和下地壳 (薛怀民等, 2010)

Fig. 8 The diagram of Sr-Nd isotopic of the concealed granite in Donggushan area

Data sources: EM I and EM II mantle component (Zindler and Hart, 1986); DM, OIB and mantle sequence (after Yang et al., 2005); The upper and lower crust of the North China Block (Jahn et al., 1999); The upper and lower Yangtze crust (after Xue et al., 2010)

5 讨论

5.1 成岩年代格架

本文通过 LA-ICP MS 定年技术对东顾山隐伏花岗岩进行了锆石 U-Pb 定年研究, 获得成岩年龄值为 $96.7 \pm 1.3 \text{Ma}$, 该年龄与聂利青等 (2016) 测得的 LA-ICP MS 锆石 U-Pb 年龄 ($99.9 \pm 1.7 \text{Ma} \sim 99.7 \pm 1.5 \text{Ma}$) 在误差范围内一致, 代表了岩浆冷凝固结的时间。东顾山隐伏花岗岩为晚白垩世早期岩浆活动的产物。

长期以来, 长江中下游成矿带的岩浆活动时空格架一直是人们所关注的问题。前人已对成矿带的岩浆活动时限进行了大量的年代学研究。周涛发等 (2008) 将长江中下游地区岩浆岩成岩时代大致划分为 $145 \sim 136 \text{Ma}$ 、 $135 \sim 127 \text{Ma}$ 、

表4 东顾山隐伏花岗岩 LA-ICP MS 锆石 U-Pb 分析结果

Table 4 LA-ICP MS zircons U-Pb isotopic date of concealed granite intrusion in Donggushan area

Spot No.	Contents (× 10 ⁻⁶)				²⁰⁷ Pb/ ²⁰⁶ Pb		²⁰⁷ Pb/ ²³⁵ U		²⁰⁶ Pb/ ²³⁸ U		²⁰⁷ Pb/ ²³⁵ U		²⁰⁶ Pb/ ²³⁸ U	
	Pb	Th	U	Th/U	Ratio	1σ	Ratio	1σ	Ratio	1σ	Age (Ma)	1σ	Age (Ma)	1σ
-01	11.66	209	154	1.36	0.0532	0.0034	0.1077	0.0061	0.0147	0.0003	103.8	5.6	93.9	1.6
-03	11.37	216	149	1.45	0.0530	0.0028	0.1061	0.0048	0.0145	0.0002	102.4	4.4	92.9	1.5
-05	10.56	220	136	1.62	0.0509	0.0030	0.1042	0.0054	0.0148	0.0003	100.6	5.0	95.0	1.6
-08	6.23	97	80	1.21	0.0494	0.0037	0.1041	0.0070	0.0153	0.0003	100.5	6.5	97.7	1.8
-09	8.83	127	96	1.33	0.0535	0.0049	0.1107	0.0095	0.0150	0.0004	106.6	8.7	96.1	2.2
-10	12.13	182	140	1.30	0.0499	0.0039	0.1046	0.0074	0.0152	0.0003	101.1	6.8	97.3	1.9
-12	12.04	221	160	1.38	0.0467	0.0032	0.0946	0.0058	0.0147	0.0003	91.8	5.4	94.1	1.6
-18	12.90	283	159	1.78	0.0508	0.0031	0.1042	0.0055	0.0149	0.0003	100.6	5.1	95.1	1.6
-24	19.18	255	221	1.15	0.0491	0.0028	0.1050	0.0050	0.0155	0.0003	101.4	4.6	99.1	1.6
-27	53.37	1114	626	1.78	0.0516	0.0023	0.1095	0.0038	0.0154	0.0002	105.5	3.5	98.4	1.4
-30	12.79	191	154	1.24	0.0508	0.0038	0.1083	0.0074	0.0155	0.0003	104.4	6.8	98.9	1.9
-31	20.97	297	263	1.13	0.0505	0.0024	0.1085	0.0040	0.0156	0.0002	104.6	3.7	99.7	1.5
-33	14.39	234	175	1.34	0.0516	0.0030	0.1089	0.0056	0.0153	0.0003	105.0	5.1	97.9	1.6
-36	27.54	263	366	0.72	0.0463	0.0020	0.0981	0.0033	0.0154	0.0002	95.0	3.1	98.2	1.4

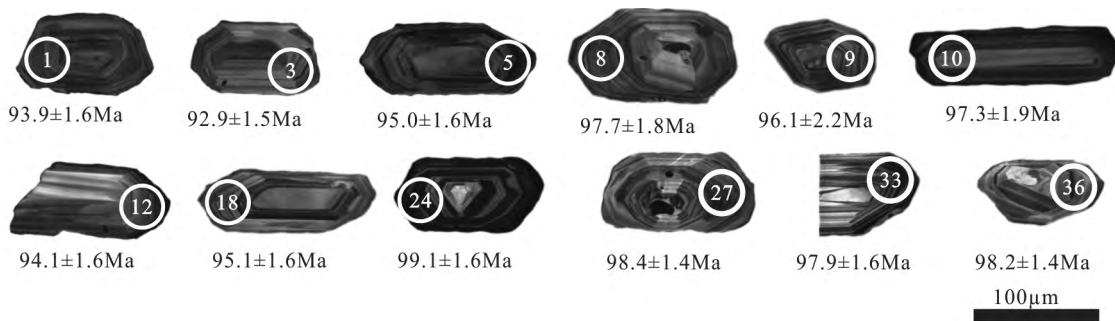


图9 东顾山隐伏花岗岩样品部分锆石阴极发光 (CL) 图像及测试位置

Fig.9 Cathodoluminescence (CL) images of selected zircons for concealed granite in Donggushan area

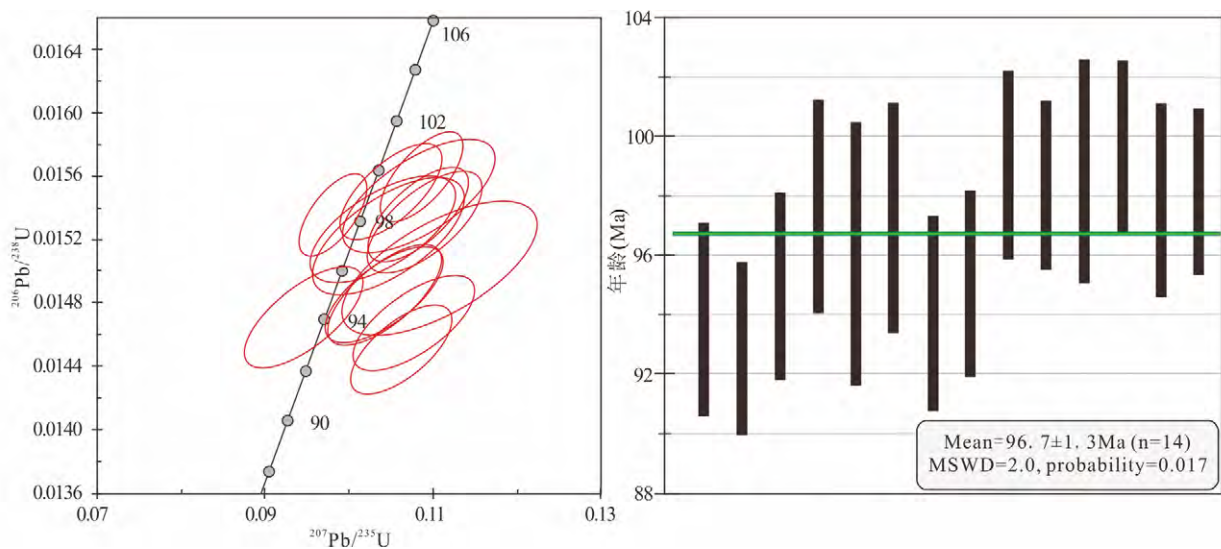


图10 东顾山隐伏花岗岩锆石 U-Pb 年龄谱和图

Fig.10 Zircon U-Pb concordia diagram of concealed granite in Donggushan area

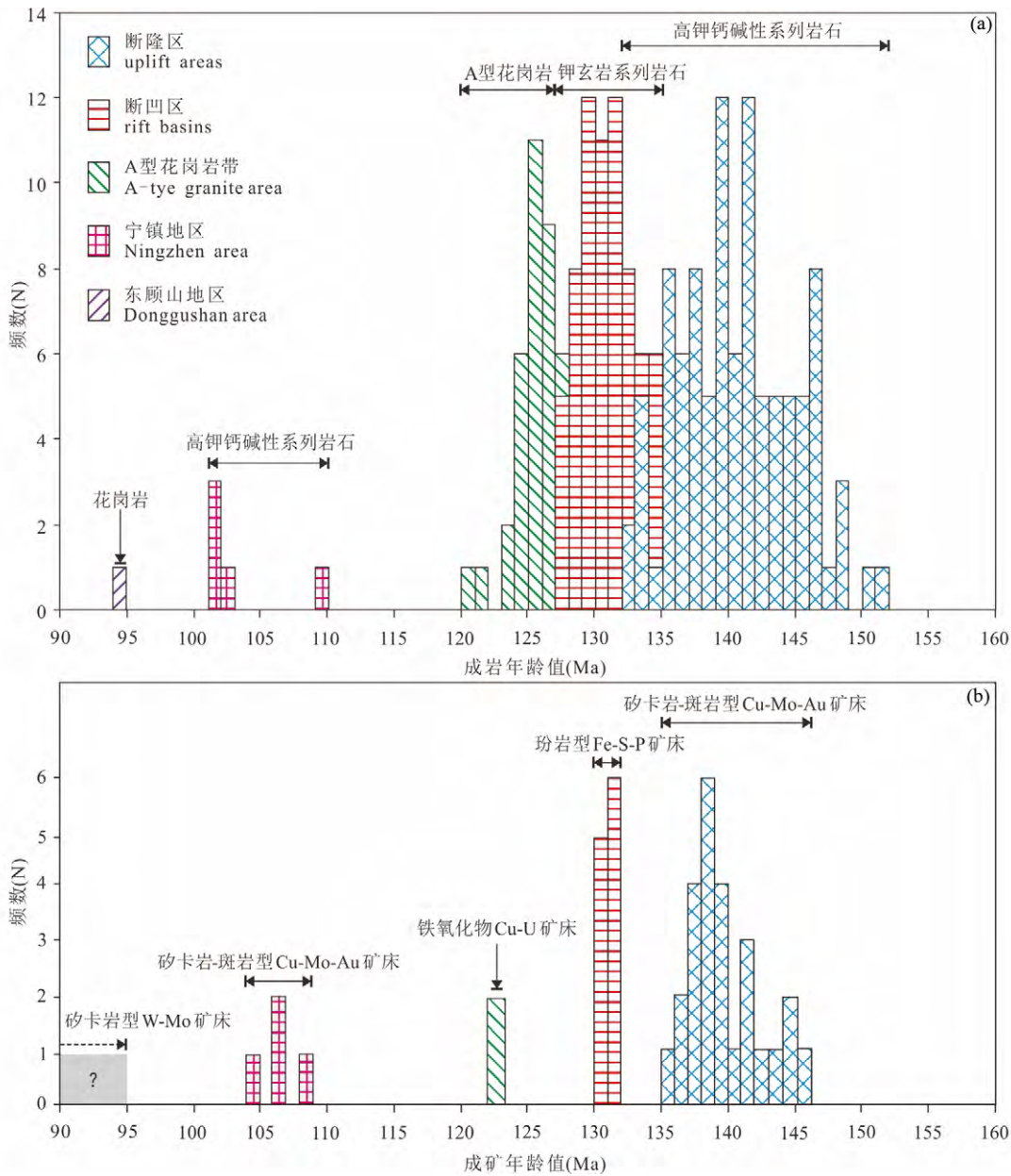


图 11 长江中下游地区岩浆活动 (a) 与成矿作用 (b) 时限直方图

主要成岩年龄数据来源: 鄂东南地区(周珣若等,1993;薛怀民等,2006;Li *et al.*,2009,2010;Xie *et al.*,2011,2012);九瑞地区(Li *et al.*,2010;陈志洪等,2011;Yang *et al.*,2011;Xie *et al.*,2011);安庆-贵池地区(陈江峰等,1991,2005;李波等,2004;Wang *et al.*,2006;Wu *et al.*,2012);铜陵地区(周泰禧等,1987;吴才来等,1996;王彦斌等,2004a,b,c;徐夕生等,2004;杨小男等,2008;谢建成等,2008;徐晓春等,2008;吴淦国等,2008);庐枞地区(Wang and McDougall,1980;刘洪等,2002;周涛发等,2007,2008,2010;薛怀民等,2010;曾键年等,2010;张乐骏,2011;范裕等,2014);宁芜地区(Wang and McDougall,1980;岳文浙和丁保良,1999;张旗等,2003;Yan *et al.*,2009;薛怀民等,2010;范裕等,2010;周涛发等,2011;袁峰等,2011);宁镇地区(徐祥和邢凤鸣,1994;孙洋等,2014;曾键年等,2013;刘建敏等,2014);滁州盆地(谢成龙等,2007;马芳和薛怀民,2011). 主要成矿年龄数据来源: 断隆区(毛景文等,2004;Zhou *et al.*,2007;王彦斌等,2004d;张乐骏等,2008;蒋少涌等,2010);断凹区(Zhou *et al.*,2011;张乐骏,2011;周涛发等,2011;张乐骏等,2010;范裕等,2014);A型花岗岩带(范裕等,2008;Mao *et al.*,2011);宁镇地区(王立本等,1997;孙洋等,2014;关俊朋等,2015)

Fig. 11 Age histogram of magmatism (a) and mineralization (b) in the Middle and Lower Reaches of Yangtze River

Data sources of main diagenetic age: Edongnan district (Zhou *et al.*,1994;Xue *et al.*,2006;Li *et al.*,2009,2010;Xie *et al.*,2011,2012); Jiurui district (Li *et al.*,2010;Chen *et al.*,2011;Yang *et al.*,2011;Xie *et al.*,2011); Anqing-Guichi district (Chen *et al.*,1991,2005;Li *et al.*,2004;Wang *et al.*,2006;Wu *et al.*,2012); Tongling district (Zhou *et al.*,1987;Wu *et al.*,1996;Wang *et al.*,2004a,b,c;Xu *et al.*,2004;Yang *et al.*,2008;Xie *et al.*,2008;Xu *et al.*,2008;Wu *et al.*,2008); Luzong district (Wang and McDougall,1980;Liu *et al.*,2002;

Zhou *et al.*, 2007, 2008, 2010; Xue *et al.*, 2010; Zeng *et al.*, 2010; Zhang *et al.*, 2011; Fan *et al.*, 2014; Ningwu district (Wang and McDougall, 1980; Yue and Ding, 1999; Zhang *et al.*, 2003; Yan *et al.*, 2009; Xue *et al.*, 2010; Fan *et al.*, 2010; Zhou *et al.*, 2011; Yuan *et al.*, 2011); Ningzhen district (Xu and Xing, 1994; Sun *et al.*, 2014; Zeng *et al.*, 2013; Liu *et al.*, 2014); Chuzhou volcanic basin (Xie *et al.*, 2007; Ma and Xue, 2011). Data sources of main metallogenic age: uplift areas (Mao *et al.*, 2004; Zhou *et al.*, 2007; Wang *et al.*, 2004d; Zhang *et al.*, 2008; Jiang *et al.*, 2010); rift basins (Zhou *et al.*, 2011; Zhang, 2011; Zhou *et al.*, 2011; Zhang *et al.*, 2010; Fan *et al.*, 2014); A-type granite zone (Fan *et al.*, 2008; Mao *et al.*, 2011); Ningzhen area (Wang *et al.*, 1997; Sun *et al.*, 2014; Guan *et al.*, 2015)

126 ~ 123Ma 三个时期, 认为 145 ~ 136Ma 的岩浆活动主要发生在断隆区(如九瑞、铜陵、贵池-安庆等); 135 ~ 127Ma 的岩浆活动主要发育在断凹区(如庐枞盆地、宁芜盆地等), 之后的 A 型花岗岩集中形成于 126 ~ 123Ma, 既可以产于断隆区, 又可以产出于断凹区。近年来, 随着研究的不断深入, 有学者对宁镇矿集区(断隆区)的岩浆岩成岩时序进行了重新厘定, 提出宁镇地区岩浆活动为长江中下游成矿带大规模岩浆活动的新一期成岩事件(曾键年等, 2013; 孙洋等, 2014; 王小龙等, 2014; 刘建敏等, 2014; 关俊朋等, 2015), 并将长江中下游地区中生代岩浆活动时限重新划分为四期(152 ~ 135Ma, 135 ~ 127Ma, 127 ~ 121Ma, 109 ~ 101Ma(孙洋等, 2014); 148 ~ 133Ma, 131 ~ 127Ma, 126 ~ 123Ma, 109 ~ 102Ma(刘建敏等, 2014))。

相比断隆区与断凹区岩浆活动时限差(7 ~ 13Myr)、断凹区与 A 型花岗岩岩浆活动时限差(3 ~ 5Myr)及 A 型花岗岩与宁镇地区岩浆活动时限差(20Myr 左右), 东顾山隐伏花岗岩的形成时代与宁镇地区岩浆活动时限相差近 8 ~ 15Myr。在成岩时代上, 东顾山隐伏花岗岩不可能与宁镇地区岩浆岩以区域上的同期岩浆活动来解释, 而可能代表了长江中下游成矿带大规模岩浆活动中新一期成岩事件。对此, 结合近年来发表的长江中下游地区高精度的岩浆岩年龄数据(包括 SHRIMP 锆石 U-Pb 法、LA-ICP MS 锆石 U-Pb 法, 年龄数据略), 通过梳理和对比, 本文重新构筑了长江中下游成矿带中生代岩浆岩年代学格架, 共划分为五个成岩作用阶段(图 11a)。第一阶段: 148 ~ 135Ma, 以高钾钙碱性系列侵入岩为主, 主要发育在断隆区的九瑞、安庆-贵池及铜陵地区; 第二阶段: 135 ~ 127Ma, 以橄榄安粗岩系次火山-火山岩为主, 主要发育在宁芜盆地、庐枞盆地及滁州盆地内; 第三阶段: 127 ~ 121Ma, 以 A 型花岗岩为主, 主要发育在沿江两岸; 第四阶段: 109 ~ 102Ma, 以高钾钙碱性系列岩浆岩为主, 主要发育在断隆区的宁镇地区; 第五阶段岩浆活动时限为 $96\text{Ma} \pm$, 以东顾山隐伏花岗岩侵入体为代表。

5.2 岩石成因

5.2.1 岩石种属

断隆区和宁镇地区岩浆岩相对富碱、富钾, 以高钾钙碱性系列为主, 且具有埃达克岩相似的地球化学特征; 断凹区岩浆岩具有高钾、富碱的特征, 属于橄榄安粗岩系列岩石。东顾山隐伏花岗岩组合指数 $\sigma = 2.06 \sim 2.76 (< 3.3)$, 属于钙碱性系列, 在 $\text{SiO}_2\text{-K}_2\text{O}$ 关系中, 表现为高钾特征, 属于高钾钙碱性系列岩石, 与断隆区岩石系列一致。

研究认为, 埃达克质岩石具有高 $\text{Sr} (> 400 \times 10^{-6})$ 、强烈

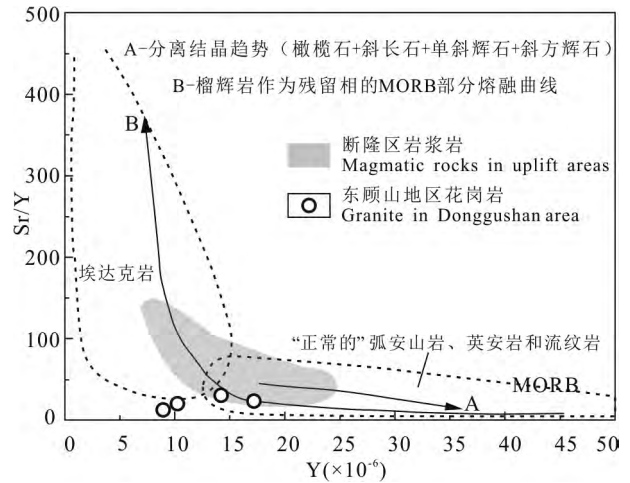


图 12 东顾山隐伏花岗岩 Sr/Y-Y 图解

Fig. 12 Sr/Y vs. Y diagram of concealed granite in Donggushan area

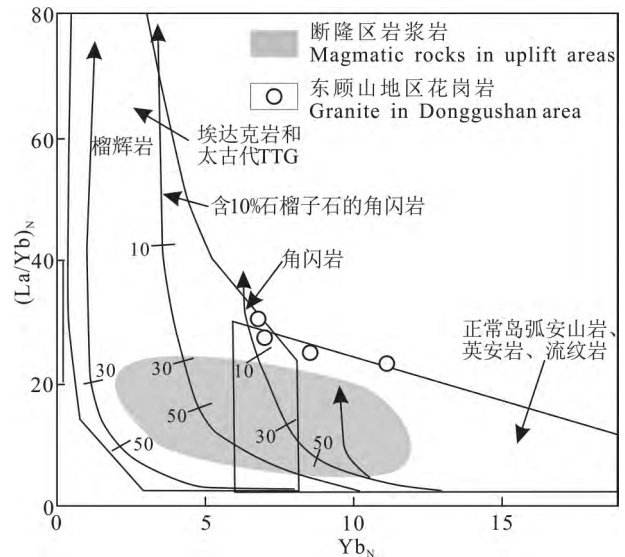


图 13 东顾山隐伏花岗岩 $(\text{La}/\text{Yb})_N$ - Yb_N 图解

Fig. 13 $(\text{La}/\text{Yb})_N$ vs. Yb_N diagram of concealed granite in Donggushan area

亏损 HREE 和 Y ($Y < 18 \times 10^{-6}$, $\text{Yb} < 1.9 \times 10^{-6}$)、高 Sr/Y (> 20)、 $\text{La}/\text{Yb} > 20$ 、 $\text{Pb}/\text{Nd} = 0.329 \sim 0.390$ 及 Eu 正异常或弱负异常等特征(张旗等, 2001, 2003, 2009; 汪洋等, 2004; 王强等, 2008)。本区隐伏花岗岩 $\text{Sr} = 146 \times 10^{-6} \sim 460 \times 10^{-6}$, 平均值为 323×10^{-6} , $Y = 8.73 \sim 17.1$, 平均 12.5, $\text{Sr}/Y = 16.7 \sim 32.3$, 平均值为 24.7, $\text{Yb} = 1.15 \sim 1.89$, $\text{La}/\text{Yb} =$

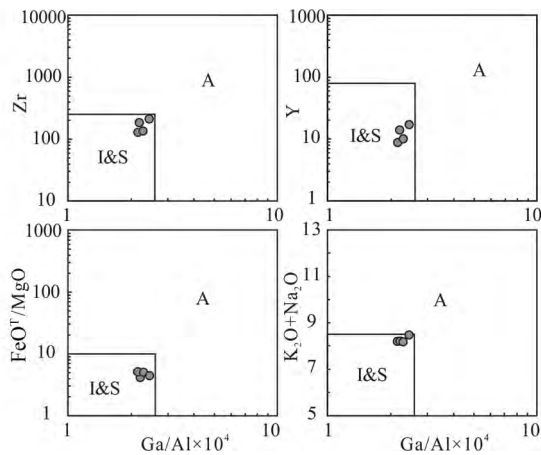


图 14 东顾山隐伏花岗岩判别图解

Fig. 14 Discrimination diagram of concealed granite in Donggushan area

37.2 ~ 44.7, $Pb/Nd = 0.392 \sim 1.172$, 在 $Sr/Y-Y$ 图解(图 12) 和 $(La/Yb)_N-Yb_N$ 图解(图 13) 中, 样品点主要分布在经典岛弧岩石范围内及附近, 说明东顾山隐伏花岗岩不具有埃达克质岩石的特征, 有别于断隆区岩浆岩和宁镇地区岩浆岩。

在岩石化学成分上, A 型花岗岩高硅、富碱、贫钙、镁, 具有高的 $(K_2O + Na_2O)/Al_2O_3$ 和 FeO^T/MgO 比值, 富 Rb、Th、Nb、Ta、Zr、Hf、Ga、Y, 贫 Ba、Sr、Ti、P、Cr、Co、Ni、V, Ga/Al 值高 (Collins *et al.*, 1982; Whalen *et al.*, 1987)。稀土元素含量高, 常为轻稀土元素富集型并且其配分模式呈海鸥型展布 (杨富贵等, 1999; Marks *et al.*, 2004)。张旗等 (2012) 认为稀土元素配分模式和微量元素蛛网图的联用是判别 A 型花岗岩最为有效的方法, 即强烈亏损 Sr、Ba、Ti、P, 强烈亏损 Eu。东顾山地区隐伏花岗岩与沿江典型 A 型花岗岩相比, Sr、Ba、Ti、P 相对富集, 也不具有强烈的 Eu 负异常, 且 FeO^T/MgO 和 Ga/Al 比值低, 亏损 Nb、Zr、Hf、Y, 在 10000Ga/Al 系列判断图解(图 14) 中, 所有样品均位于非 A 型花岗岩区域。东顾山隐伏花岗岩缺少 S 型花岗岩的特征的富铝矿物, 且铝饱和指数 ANCK 的值都小于 1, 表明其具有 I 型花岗岩的特征。实验岩石学表明, 弱过铝质岩浆中 P_2O_5 随着 SiO_2 的增加呈负相关的特征, 也成功地区分了 I 型和 S 型花岗岩 (Wu *et al.*, 2003; Li *et al.*, 2007), 本区花岗岩的 P_2O_5 随着 SiO_2 的增加而减少, 表明东顾山地区隐伏花岗岩为弱过铝质高钾钙碱性系列 I 型花岗岩。

5.2.2 岩浆源区

目前, 对于断隆区高钾钙碱性系列岩石的岩浆源区的研究具有较大争议, 主要观点包括幔源岩浆和壳源岩浆的混合 (邢凤鸣和徐祥, 1996; 吴才来等, 2003; 王强等, 2003)、热的玄武岩浆地底侵引起加厚和/或拆沉下地壳的部分熔融 (邢凤鸣和徐祥, 1999; Li *et al.*, 2009; 徐晓春等, 2012) 以及洋脊俯冲背景下洋壳的部分熔融形成 (Ling *et al.*, 2009, 2011; 孙卫东等, 2010)。大多数学者认为断凹区橄榄安粗

岩系列岩石的岩浆起源于富集型岩石圈地幔, 并有软流圈物质的加入 (薛怀民等, 2012, 2016; Chen *et al.*, 2014; Xue *et al.*, 2015; Yan *et al.*, 2015), 而富集型岩石圈地幔的形成可能是由古太平洋板块俯冲析出的流体/熔体交代形成 (刘洪等, 2002; Wang *et al.*, 2006; Li *et al.*, 2013), 或是由新元古代扬子地块周边的俯冲交代形成 (Tang *et al.*, 2012; Chen *et al.*, 2014)。对于沿江 A 型花岗岩, 岩浆起源于岩石圈地幔, 并经历不同程度的地壳物质的混染 (邢凤鸣和徐祥, 1999; 曹毅等, 2008; 向文帅等, 2009; Wu *et al.*, 2012), 或为中-新元古代增生地壳的深熔作用形成 (闫峻等, 2012; 彭戈等, 2012)。宁镇地区中生代中酸性岩浆岩具有埃达克质岩石的地球化学特征, 岩浆来源于加厚下地壳的部分熔融 (许继峰等, 2001; Wang *et al.*, 2014), 也有学者认为其属于壳幔型岩浆成因 (宁仁祖和陈根生, 1989; 徐莺, 2010)。

东顾山地区隐伏花岗岩富碱、高钾、贫铁、贫钙、富集轻稀土元素, 亏损重稀土元素, Rb、Th、U 富集, Ba、Sr 偏低, 亏损高场强元素 Nb、Ti、P 等, La/Nb 值变化范围小, 平均值为 2.14 (与大陆地壳 La/Nb 比值 2.2) (黎彤等, 1999) 一致, 具有典型的壳源花岗岩的特征。源于软流圈地幔的岩浆以高的 $\epsilon_{Nd}(t)$ (约为 +8)、 $(^{87}Sr/^{86}Sr)_i$ (约为 0.703) 为特征, 源于岩石圈地幔的岩浆的 $\epsilon_{Nd}(t) \approx -9 \pm 2$ 、 $(^{87}Sr/^{86}Sr)_i = 0.7073 \sim 0.7097$ (DePaolo and Daley, 2000)。在 Sr-Nd 同位素组成上, 东顾山隐伏花岗岩 ($\epsilon_{Nd}(t) = -18.2 \sim -17.1$ ($(^{87}Sr/^{86}Sr)_i = 0.7060 \sim 0.7063$) 与断隆区、断凹区、宁镇地区岩浆岩及 A 型花岗岩存在较大的差异, 且具有一定的特殊性, 但与邻区洪镇花岗岩 ($\epsilon_{Nd}(t) = -23.6 \sim -11.6$ ($(^{87}Sr/^{86}Sr)_i = 0.7062 \sim 0.7083$) (Wang *et al.*, 2004; 王斌等, 2012)) 和大别山早期 (143 ~ 129Ma) 花岗岩 ($\epsilon_{Nd}(t) = -19.4 \sim -14.6$ ($(^{87}Sr/^{86}Sr)_i = 0.7067 \sim 0.7087$) (Wang *et al.*, 2007)) 相似。大别山早期花岗岩具有高 Sr 低 MgO、Y、Yb、Ni, 以及高的 Sr/Y、La/Yb 比值的特征 (Wang *et al.*, 2007; Xu *et al.*, 2012), 普遍具有埃达克岩的性质, 其形成可能与华南板块俯冲导致的地壳增厚及部分熔融有关 (Wang *et al.*, 2007; Zhao *et al.*, 2011)。洪镇花岗岩具有高 SiO_2 和 K_2O/Na_2O , 高 Sr 低 MgO、Y、Ni、V 的地球化学特征 (Wang *et al.*, 2004; 王斌等, 2012), 对于其岩浆来源及岩石成因, 可能来自迄今尚未出露的古老基底 (陈江峰等, 1993) 或董岭群古老基底的重熔 (王斌等, 2012), 或形成于古老大陆下地壳高钾基性物质的部分熔融 (Wang *et al.*, 2004), 也有人认为洪镇花岗岩与大别山早期花岗岩具有相似的源区, 均来源于大别造山带加厚的地壳 (余顶杰等, 2016)。可以看出, 东顾山隐伏花岗岩、宁镇花岗岩及大别山花岗岩均来自地壳物质的部分熔融, 无地幔物质的加入, 所不同的是东顾山花岗岩不具有埃达克质岩石的地球化学特征, 并非来自加厚下地壳的熔融。在 $\epsilon_{Nd}(t) - I_{Sr}$ 图解中 (图 8), 东顾山隐伏花岗岩靠近扬子下地壳区域, 远离扬子克拉通岩石圈地幔, 说明本区隐伏花岗岩岩浆来源于扬子下地壳的部分熔融。

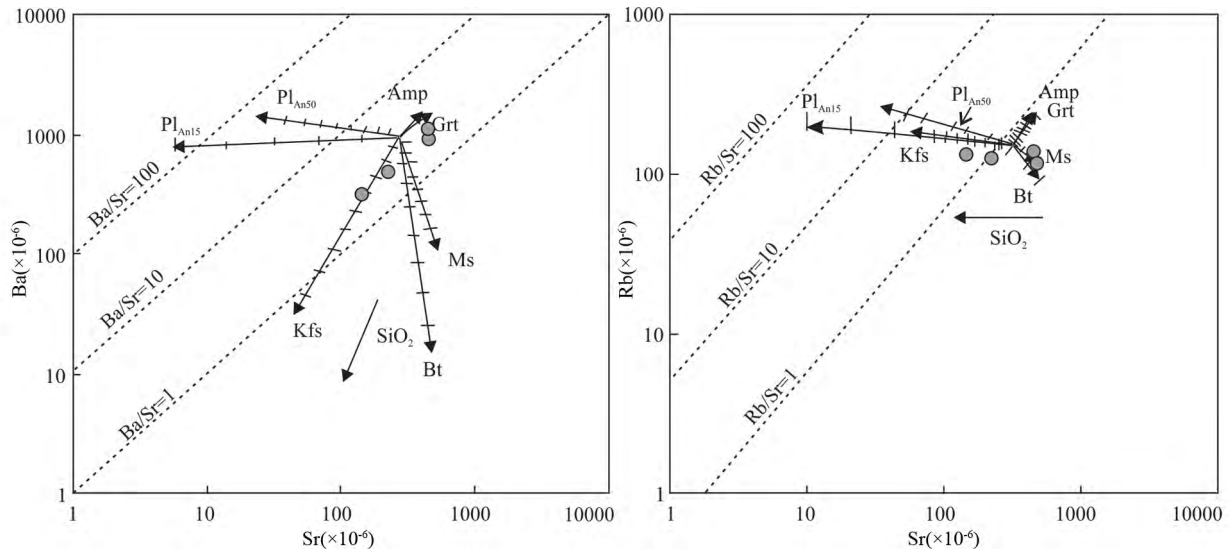


图 15 东顾山隐伏花岗岩造岩矿物结晶分异判断图解(底图据 Janoušek *et al.*, 2004)

Pl_{An50}-斜长石 (An = 50); Pl_{An15}-斜长石 (An = 15); Kfs-钾长石; Bt-黑云母; Ms-白云母; Grt-石榴子石; Amp-角闪石

Fig. 15 Ba vs. Sr and Rb vs. Sr plots for the concealed granite in Donggushan area (after Janoušek *et al.*, 2004)

Pl_{An50}-Plagioclase (An = 50); Pl_{An15}-Plagioclase (An = 15); Kfs-K-feldspar; Bt-biotite; Ms-muscovite; Grt-garnet; Amp-amphibole

花岗岩高 Si, 低 Ba、Sr, 亏损 Nb、Ta、Sr、P、Ti 和 Eu 等元素, 指示母岩浆经历了分离结晶演化 (Wu *et al.*, 2003)。在 Ba-Sr、Rb-Sr 对数图解中 (图 15) 随着 Ba 含量的升高, Sr 含量沿钾长石结晶分异演化线从 146×10^{-6} 增加到 460×10^{-6} , 这种变化趋势证明了侵入体在形成过程中经历了钾长石和斜长石的分异结晶。东顾山隐伏花岗岩 Yb、Y 含量低, 且亏损 HREE, 在 Ba-Sr 图解中, 部分样品分布在向石榴子石演化趋势线上, 说明源区可能有石榴子石的存在。但侵入体贫 Sr、低 Al, 具有 Eu 负异常, 说明残留相中也有斜长石存在。由于贫 Sr、低 Yb, 本区侵入岩属于低 Sr 低 Y 型花岗岩 (张旗等, 2008), 形成于中等压力条件下。Wang *et al.* (2004) 通过对洪镇花岗岩地球化学的研究认为, 低 Y、Yb, 高 Sr/Y、La/Yb 比值的特征表明了岩浆源区残留石榴子石, 中稀土元素的亏损及 K/Rb 比值偏低的特征暗示源区可能有角闪石残留以及低 CaO、Al₂O₃、低 Ba/Sr 值以及弱 Sr 负异常表明源区也有斜长石残留。大别山早期埃达克质花岗岩高 Al₂O₃、Sr 低 Y、Yb, 弱 Eu 正异常和 Sr 正异常的特征表明源区残留石榴子石及少量或者不含斜长石, Na/Ta 比值接近甚至高于原始地幔的值, 暗示源区残留金红石 (Wang *et al.*, 2004; He *et al.*, 2011)。

5.2.3 构造背景及动力学过程

T₃-J₂ (约 220Ma) 时期中国东部在完成华北板块与扬子板块之间的碰撞并最终焊接后, 长江中下游地区转为受太平洋构造体制制约的大陆岩石圈演化阶段 (邓晋福等, 2004; 杜杨松等, 2007; 董树文等, 2007, 2011)。晚侏罗纪 (165 ± 5Ma ~ 145Ma) 古太平洋板块开始俯冲, 造成中国东部受挤压整体抬升形成高原, 遭受侵蚀而缺失晚侏罗世的沉积, 岩石

圈迅速增厚 (董树文等, 1993; 张旗等, 2001); 随后进入碰撞造山后的应力转换期 (145 ~ 136Ma), 构造应力由挤压向拉伸过渡, 岩石圈地幔发生部分熔融, 形成了埃达克质母岩浆 (王强等, 2004; 蒋少涌等, 2008), 岩浆在上升过程中不同程度地受到地壳物质的混染, 经过混合的岩浆经过复杂的演化过程后, 侵位到地壳浅部, 形成了以中酸性为主的具埃达克质特征的侵入体及砂卡岩-斑岩型铜-金成矿作用 (毛景文等, 2005; 谢桂青等, 2006; 王建中等, 2008; 徐晓春等, 2008, 2012; 曾键年等, 2010)。这一过程主要形成于断隆区。约 135 ~ 127Ma, 长江中下游地区进入拉伸构造期, 岩石圈拆沉、减薄, 软流圈上涌, 富集型的岩石圈地幔发生部分熔融, 形成的岩浆在壳幔过渡带附近聚集发生分离结晶作用 (薛怀民等, 2016)。随后岩浆上升到地壳浅部, 发生的火山作用形成了一系列断陷火山岩盆地和以橄榄安岩系列为主的火山作用及与火山-次火山岩有关的玢岩型铁硫磷成矿作用 (任启江等, 1991; 周涛发等, 2010; 薛怀民等, 2012)。这一过程主要发生在断凹区。126 ~ 123Ma, 拉伸作用进一步加强, 岩石圈减薄和伸展达到高峰期 (范裕等, 2008), 富集岩石圈地幔进一步部分熔融形成的岩浆和部分软流圈地幔减压熔融所形成的岩浆混合, 这种混合岩浆首先在地壳深部发生了镁铁(钛)质矿物的分离结晶作用, 在地壳浅部又发生了以斜长石为主, 晚期还有钾长石和黑云母的分异结晶作用, 形成了以正长岩-石英正长岩-正长花岗岩为组合的 A 型花岗岩及铁氧化物-铜-铀成矿作用 (陈一峰, 1994; 陈一峰等, 1996; 范裕等, 2008; 周涛发等, 2012; 薛怀民等, 2016), 主要分布在长江沿岸。至此, 长江中下游地区晚中生代强烈的岩浆-构造-成矿作用相对减弱。

约 110 ~ 100Ma, 太平洋板块转变 75°, 向 SW 方向俯冲,

包括长江中下游地区在内的中国东部又开始伸展活动(Sun *et al.*, 2007)。岩石圈减薄、拆沉,在宁镇地区形成了具有埃达克质岩特征的中酸性侵入体及与之有关的矽卡岩型铁矿和矽卡岩型-斑岩型铜矿成矿作用(曾键年等, 2013; 王小龙等, 2014; 孙洋等, 2014; Wang *et al.*, 2014; Xue *et al.*, 2015)。目前,对于花岗岩构造背景的研究存在较大的分歧(张旗等, 2007; 吴福元等, 2007)。肖庆辉等(2007)认为高钾钙碱性系列花岗岩可以出现在各种不同的地球动力学环境中,它们实际上指示的是一种构造体制的变化而不是一个特定的地球动力学环境。受太平洋板块的俯冲作用,下扬子地区在燕山期经历了四幕构造变形,最晚一幕逆冲推覆构造变形发生在 K_1-K_2 (100Ma) 时期,随后进入了造山后伸展崩塌,形成断陷盆地(邓晋福和吴宗絮, 2001)。根据成岩时代,东顾山地区的燕山期高钾钙碱性岩浆可能形成于由挤压向拉张过渡的构造背景下,是不同于断隆区和宁镇地区的新期岩浆作用的产物。

5.3 与岩浆活动有关的成矿作用

前人对长江中下游成矿带的成矿作用研究取得了丰硕的成果,系统划分了区内成矿系列:(1)与高钾钙碱性岩浆活动有关的矽卡岩-斑岩型 Cu-Mo-Au 矿床,以断隆区(铜陵、九瑞、安庆-贵池)及宁镇地区两期成矿为代表(毛景文等, 2004, 2009; 吴淦国等, 2008),成矿时代分别为 146~135Ma 和 108~104Ma;(2)与橄榄安粗岩有关的玢岩型 Fe-S-P 矿床,以断凹区(宁芜、庐枞)为代表(宁芜研究项目编写小组, 1978; 任启江等, 1991; 周涛发等, 2010),集中形成于 130Ma 左右;(3)与 A 型花岗岩有关的铁氧化物-铜-钼矿床(陈一峰, 1994; 陈一峰等, 1996; 郑永飞等, 1995, 1997; 周涛发等, 2012),主要形成时代为 127~123Ma(图 11b)。东顾山地区深部钨钼多金属矿化与该区隐伏花岗岩在时空和成因上有密切的关系,形成于 97Ma 左右(聂利青等, 2016),可能代表了长江中下游成矿带最晚一期成矿作用事件,但在成矿系列上,与断隆区和宁镇地区相似,同属于与高钾钙碱性岩浆活动有关的矽卡岩-斑岩型矿床。

长江中下游成矿带中已发现的具有规模的钨矿床数量较少,且均位于长江深大断裂以南,受不同时期岩浆作用的控制,如鄂东南矿集区南部阮家湾大型层控矽卡岩型 W-Cu-Mo 矿床、中南部傅家山和大冶龙角山中型矽卡岩型 W-Cu-Mo 矿床(朱增青, 1987; 舒全安等, 1992)、宁镇矿集区谏壁中型斑岩型 W-Mo 矿床(杨松生等, 1985; 马春和王素娟, 2003)及铜陵矿集区姚家岭大型锌多金属矿床中的白钨矿化(钟国雄等, 2014)。而具有大中型规模的矽卡岩-斑岩型钨钼矿床主要集中分布在长江中下游成矿带以南的皖南地区,如祁门东源斑岩型 W-Mo 矿床、青阳百丈岩矽卡岩-斑岩型 W-Mo 矿床、池州鸡头山矽卡岩-斑岩型 W-Mo 矿床等。东顾山地区与隐伏花岗岩有关的钨钼多金属矿化为长江以北地区首次发现,这一发现改变了“南钨北移”的界线不会超过

长江深大断裂带的认识(马振东等, 1999; 王登红等, 2012)指示了长江深断裂以北地区具有寻找矽卡岩-斑岩型钨钼多金属矿床的潜力。

6 结论

(1) 庐枞矿集区与钨矿床有关的花岗岩属于弱过铝质高钾钙碱性系列 I 型花岗岩,岩浆来源于扬子下地壳的部分熔融,岩浆演化过程中经历了一定程度的分异结晶作用,形成于由挤压向拉张过渡的构造背景之下;

(2) 庐枞矿集区东顾山隐伏花岗岩形成于 96.7 ± 1.3 Ma,为晚白垩世早期岩浆活动的产物,其在岩石地球化学和成岩年龄上不同于断隆区高钾钙碱性系列岩石、断凹区橄榄安粗岩系列岩石、沿江 A 型花岗岩及宁镇高钾钙碱性系列岩石,可能代表了长江中下游成矿带新一期成岩成矿事件;

(3) 庐枞矿集区东顾山钨钼多金属矿化的首次发现,改变了“南钨北移”界线不超过长江深大断裂带的认识,指示长江深断裂带以北地区具有寻找矽卡岩-斑岩型钨钼多金属矿床的潜力。

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